# Principal Dynamical Components

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\*Joint work with Esteban G. Tabak, Courant Institute

1. Principal component analysis (PCA) and Autorregresive models (AR(p))

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3. Application to the Global Sea-Surface Temperature Field

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### A COUPLE OF EXAMPLES

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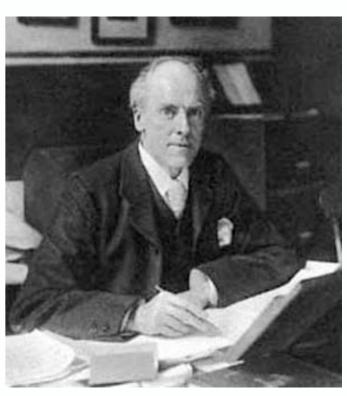
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III. On Lines and Planes of Closest Fit to Systems of Points in Space. By KARL PEARSON, F.R.S., University College, London \*.

(1) In many physical, statistical, and biological investigations it is desirable to represent a system of points in plane, three, or higher dimensioned space by the best-fitting" straight line or plane. Analytically this consists in taking



1901

#### HOW TO OBTAIN THE PC'S

#### Singular value decomposition

Given a data set  $z_1, z_2, \ldots, z_N \in \mathbb{R}^n$  (already subtracted the mean value), the first m PCs are given by  $x_j = Q'_x z_j$  where  $Q_x \in \mathbb{R}^{n \times m}$  has orthogonal columns such that the predictive uncertainty

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The matrix  $Q_x$  consists of the first m columns of U in the singular value decomposition

$$Z' = USV'$$

where  $U \in \mathbb{R}^{n \times n}$  and  $V \in \mathbb{R}^{N \times N}$  are orthogonal matrices and  $S \in \mathbb{R}^{n \times N}$  is diagonal with the eigenvalues of the covariance matrix Z'Z sorted in decreasing order  $(Z = [z_1 | \cdots | z_N])$ .

## AUTORREGRESIVE MODELS (AR(P))

#### One dimensional AR(p)

For a random process z the AR(p) model is defined as

$$z_j = b + \sum_{i=1}^p a_i z_{j-i} + \varepsilon_j,$$

where  $a_i, \ldots, a_p$  are the *parameters* of the model, b is a constant, and  $\varepsilon_j$  is white noise. The process is *stationary* if the roots of the polynomial  $x^p - \sum_{i=1}^p a_i x^{p-i}$  lie within the unit circle.

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 $a_i, \ldots, a_p$  can be estimated by solving the Yule-Walker equations



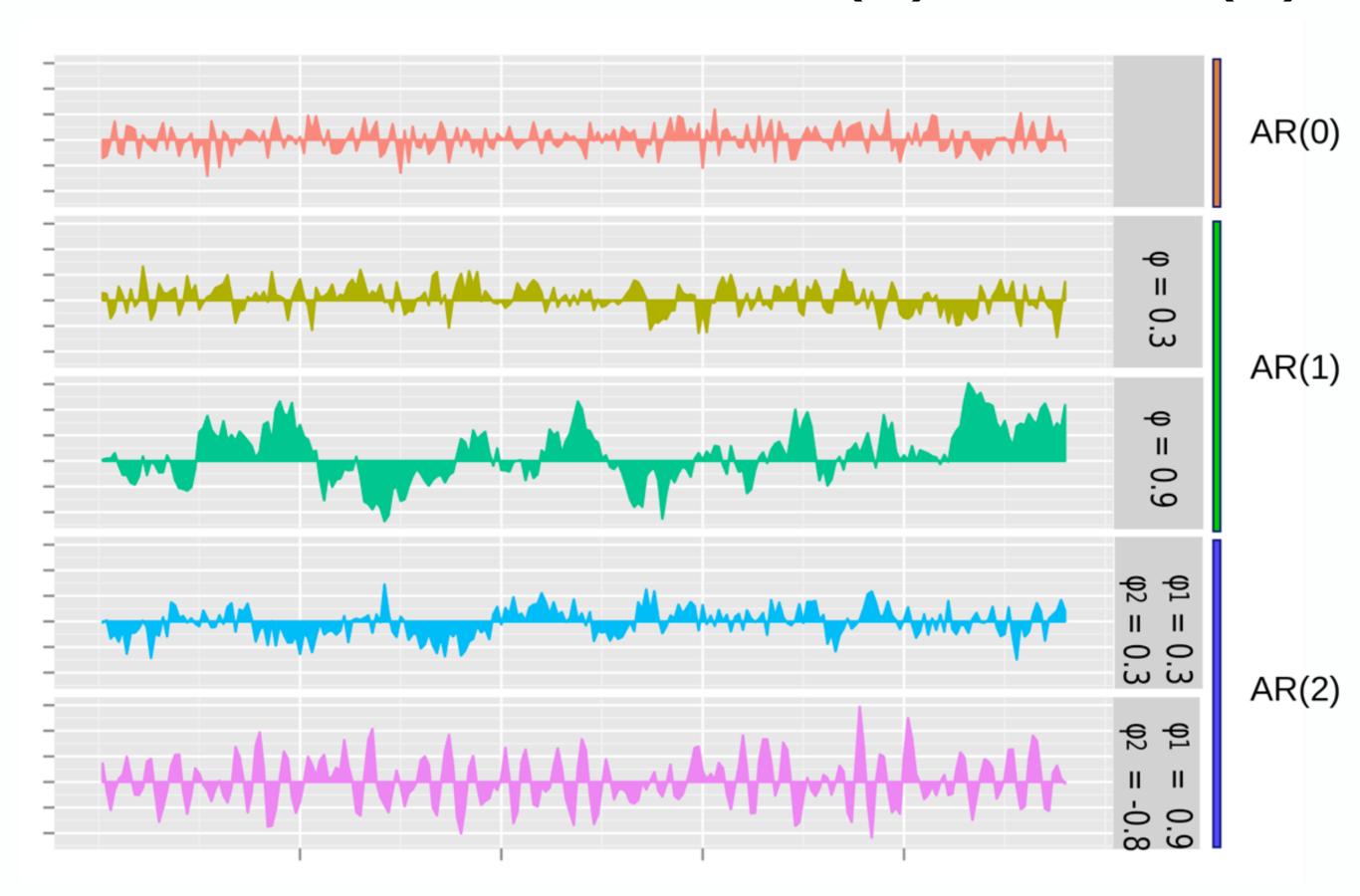
$$\begin{pmatrix} \gamma_1 \\ \gamma_2 \\ \vdots \\ \gamma_p \end{pmatrix} = \begin{pmatrix} \gamma_0 & \gamma_{-1} & \gamma_{-2} & \cdots \\ \gamma_1 & \gamma_0 & \gamma_{-1} & \cdots \\ \vdots & \vdots & \vdots & \ddots \\ \gamma_{p-1} & \gamma_{p-2} & \gamma_{p-3} & \cdots \end{pmatrix} \begin{pmatrix} a_1 \\ a_2 \\ \vdots \\ a_p \end{pmatrix},$$



1931

where  $\gamma_m = E[z_{j+m}z_j]$  are the covariance functions.

# SOME EXAMPLES OF AR(1) AND AR(2)



# VECTOR AR(P)

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Similarly an estimation of the parameters  $A_i, \ldots, A_p$  can be calculated solving the Yule-Walker equations (now block matrices).

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Given a time series  $z_j \in \mathbb{R}^n$  with transition probability  $T(z_{j+1}|z_j)$  (Markovian), PDC considers a dimensional reduction of  $T(z_{j+1}|z_j)$  in the following way

$$T(z_{j+1}|z_j) = J(z_{j+1})e(y_{j+1}|x_{j+1})d(x_{j+1}|x_j),$$

where

- $x = P_x(z(x,y)) \in \mathbb{R}^m$ ,  $y = P_y(z(x,y)) \in \mathbb{R}^{n-m}$  ( $P_x$  and  $P_y$  are projection operators) and J(z) is the Jacobian determinant of the coordinate map  $z \to (x,y)$ .
- e(y|x) is a probabilistic embedding.
- $d(x_{j+1}|x_j)$  is a reduced dynamical model.

We will focus on the case where  $P_x$  and  $P_y$  are orthogonal projections (therefore J(z) = 1) and the embedding and reduced dynamics are given by isotropic Gaussians

$$e(y|x) = \mathcal{N}(0, \sigma^2 I_{n-m}), \quad d(x_{j+1}|x_j) = \mathcal{N}(Ax_j, \sigma^2 I_m).$$

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Therefore, the log-likelihood function is given by

$$L = -\sum_{j=1}^{N-1} \left[ \frac{n}{2} \log(2\pi) + n \log(\sigma) + \frac{1}{2\sigma^2} \left( ||x_{j+1} - Ax_j||^2 + ||y_{j+1}||^2 \right) \right].$$

Maximazing L over P and A is equivalent to minimizing the cost function

$$c = \frac{1}{N-1} \sum_{j=1}^{N-1} (\|x_{j+1} - Ax_j\|^2 + \|y_{j+1}\|^2).$$

### LINEAR AND AUTONOMOUS CASE (MARKOVIAN)

For a time series  $z \in \mathbb{R}^n$  we look for a m-dimensional submanifold  $x = Q'_x z$  such that

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The minimization problem that defines  $Q = [Q_x Q_y]$  and A is

$$\min_{Q,A} c = \sum_{j=1}^{N-1} \left\| z_{j+1} - Q \begin{pmatrix} AQ_x'z_j \\ 0 \end{pmatrix} \right\|^2 = \sum_{j=1}^{N-1} \left\| \begin{pmatrix} x_{j+1} - Ax_j \\ y_{j+1} \end{pmatrix} \right\|^2.$$

Let us call  $z = \begin{pmatrix} A \\ P \end{pmatrix}$ . We look for a one-dimensional submanifold x of z

$$x = A\cos(\theta) + P\sin(\theta)$$

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$$c(\theta, a) = \sum_{j=1}^{N-1} \left\| \begin{pmatrix} A_{j+1} - \tilde{A}_{j+1} \\ P_{j+1} - \tilde{P}_{j+1} \end{pmatrix} \right\|^2 = \sum_{j=1}^{N-1} \left\| \begin{pmatrix} x_{j+1} - \tilde{x}_{j+1} \\ y_{j+1} - \tilde{y}_{j+1} \end{pmatrix} \right\|^2$$
$$= \sum_{j=1}^{N-1} \left\| \begin{pmatrix} x_{j+1} - ax_j \\ y_{j+1} \end{pmatrix} \right\|^2 = \sum_{j=1}^{N-1} (y_{j+1})^2 + (x_{j+1} - ax_j)^2$$

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By contrast, the corresponding cost function for regular principal components in this 2-dimensional scenario is

$$c_{PCA}(\theta) = \sum_{j=1}^{N} y_j^2.$$

### SYNTHETIC EXAMPLE

We created data from the dynamical model

$$x_{j+1} = ax_j + 0.3\eta_j^x$$
  
$$y_{j+1} = 0.6\eta_j^y$$

where a = 0.6, j = 1, ..., 999 and  $\eta_j^{x,y}$  are independent samples from a normal distribution.

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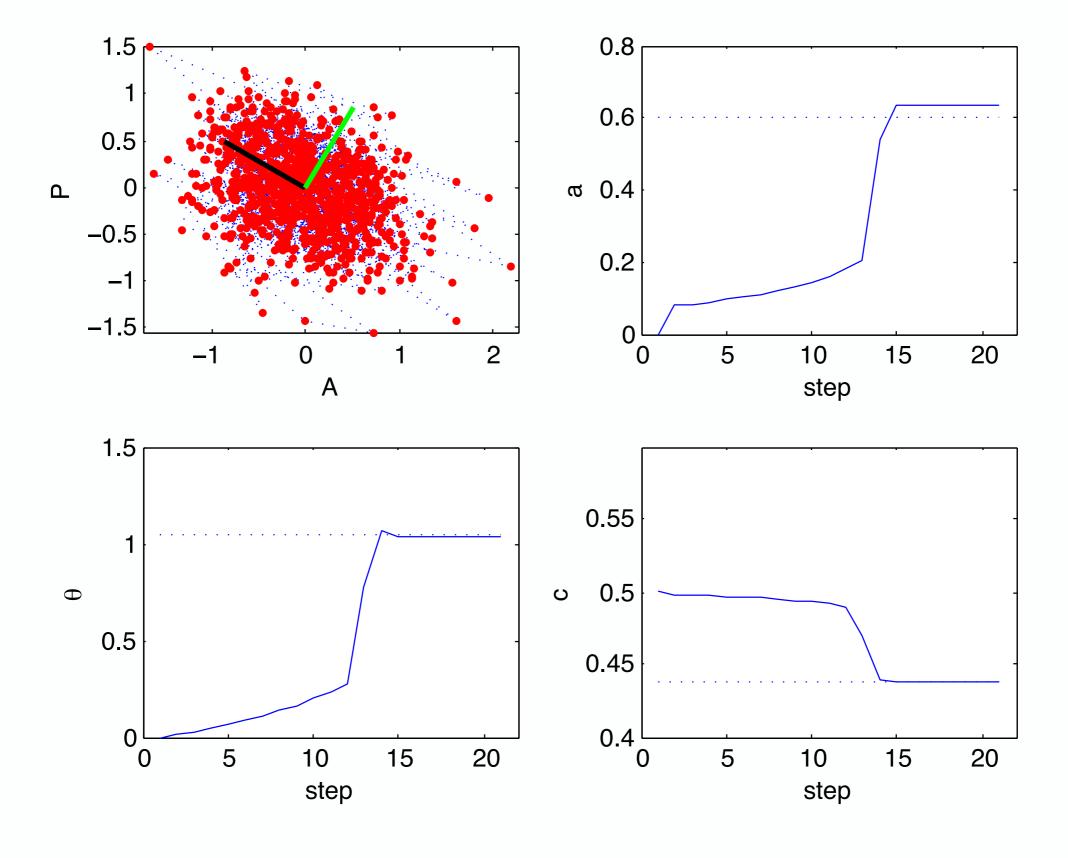
where a = 0.6, j = 1, ..., 999 and  $\eta_j^{x,y}$  are independent samples from a normal distribution.

Then we rotated the data through the angle  $\theta = \frac{\pi}{3}$ 

$$A_j = x_j \cos(\theta) - y_j \sin(\theta)$$
  
$$P_j = x_j \sin(\theta) + y_j \cos(\theta)$$

and we perform descent over the variables a and  $\theta$ .

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 $y_{j+1} = \bar{y}_{j+1} + 0.6 \eta_j^y,$ 

for j = 1, ..., 999 and we adopted the values  $a_j = \frac{6}{5}\cos^2\left(\frac{2\pi t_j}{T}\right)$  for the dynamics,  $b_j = \frac{1}{2}\sin\left(\frac{2\pi t_j}{T}\right)$  for the drift, and  $\bar{y}_j = \frac{2}{5}\cos\left(\frac{2\pi t_j}{T}\right)$  for the non-zero mean of y, where  $t_j = j$  and T = 12.

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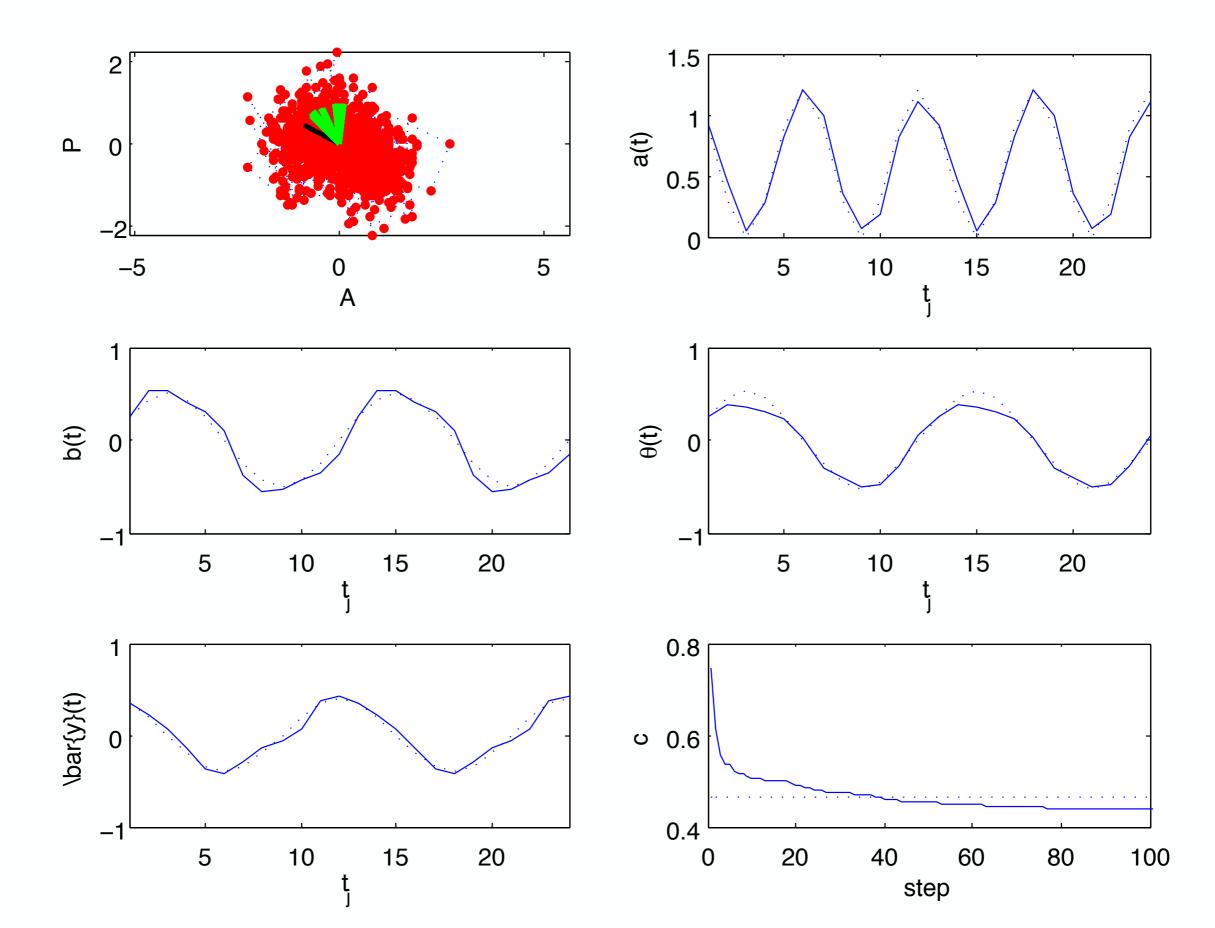
Then we rotated the data through

$$A_j = x_j \cos(\theta_j) - y_j \sin(\theta_j),$$
  

$$P_j = x_j \sin(\theta_j) + y_j \cos(\theta_j),$$

where 
$$\theta_j = \frac{\pi}{6} \sin\left(\frac{2\pi t_j}{T}\right)$$
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### NONAUTONOMOUS PROBLEMS



Take n = 2, m = 1, and r = 3, the order of the Non-Markovian process. We created data from the dynamical model

$$x_{j+1} = a_1 x_j + a_2 x_{j-1} + a_3 x_{j-2} + 0.3 \eta_j^x,$$
  
$$y_{j+1} = 0.6 \eta_j^y,$$

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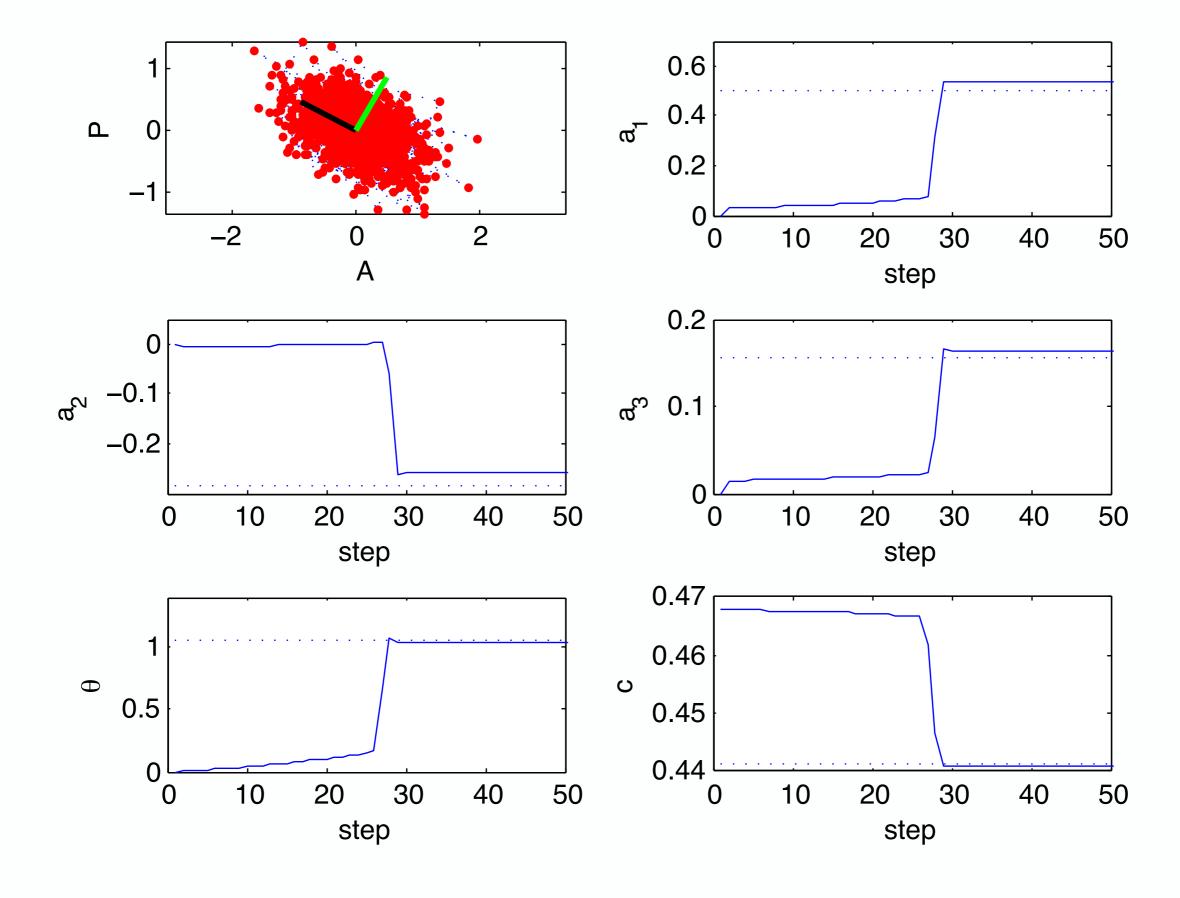
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Then, as before, we define

$$A_j = x_j \cos(\theta) - y_j \sin(\theta),$$
  

$$P_j = x_j \sin(\theta) + y_j \cos(\theta),$$

with  $\theta = \frac{\pi}{3}$ , and provide the  $A_j$  and  $P_j$  as data for the principal dynamical component routine.



#### Preliminary considerations

Database: monthly averaged extended reconstructed global sea surface temperatures based on COADS data (January 1854 to October 2009).

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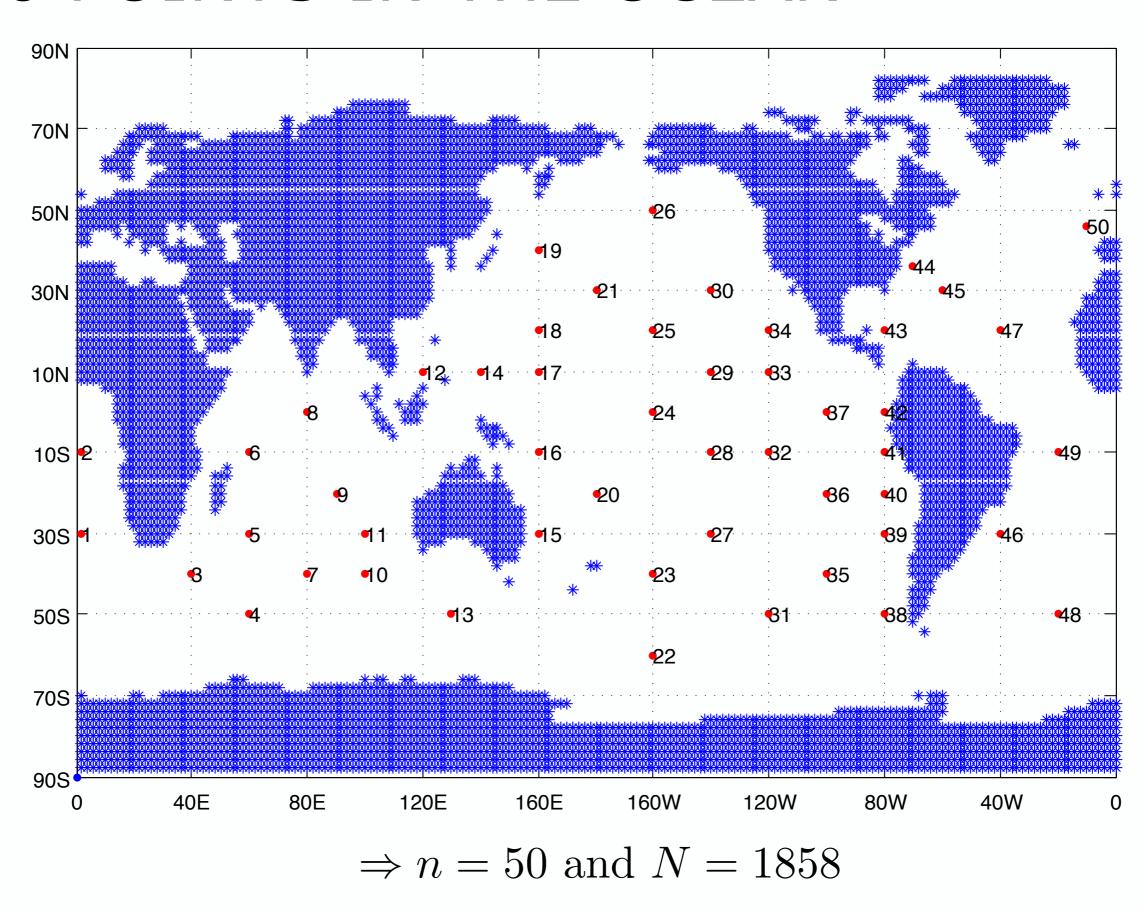
• The ocean is not an isolated player in climate dynamics: it interacts with the atmosphere and the continents, and is also affected by external conditions, like solar radiation or human-related release of  $CO_2$  into the atmosphere. The latter are examples of slowly varying external trends that fit naturally into our non-autonomous setting.

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- Even within the ocean, the surface temperature does not evolve alone: it is carried by currents, and it interacts through mixing with lower layers of the ocean. One way to account for unobserved variables is to make the model non-Markovian.

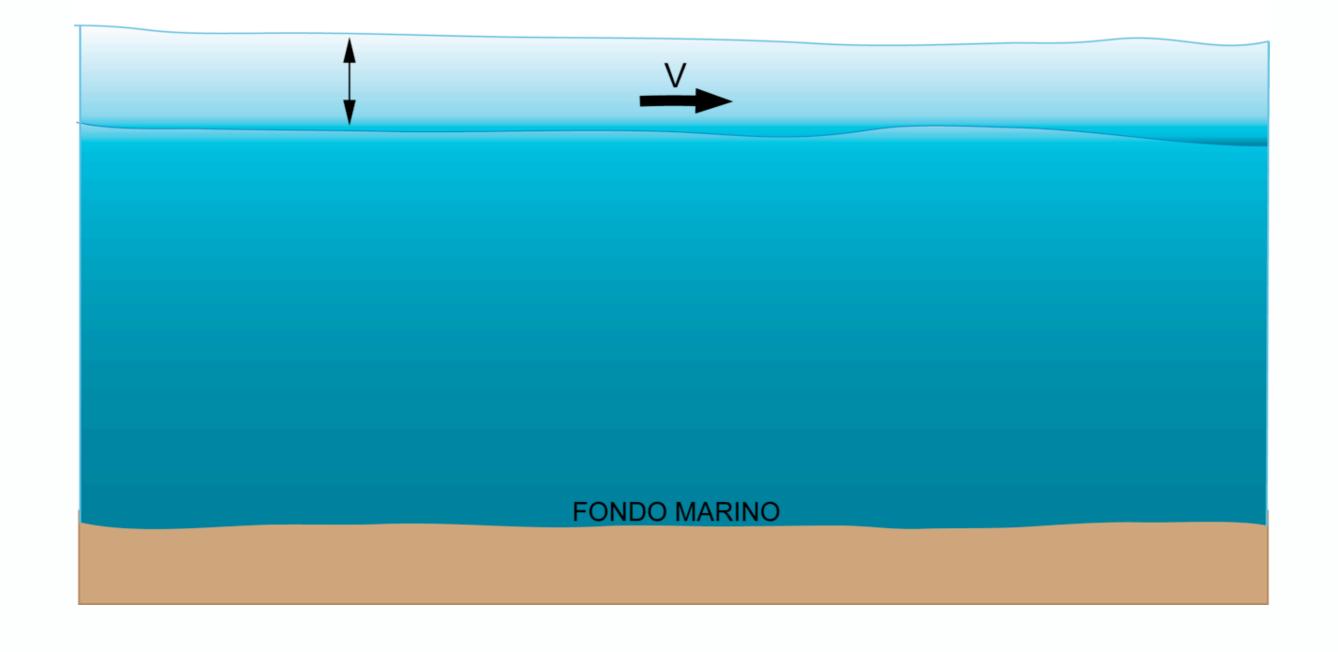
#### 50 POINTS IN THE OCEAN



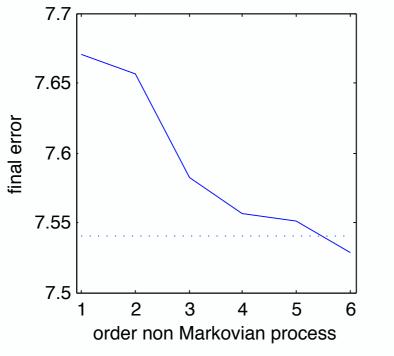
Order of the process: r = 3

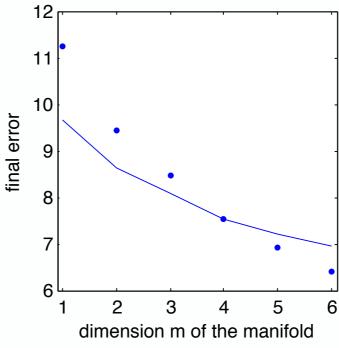
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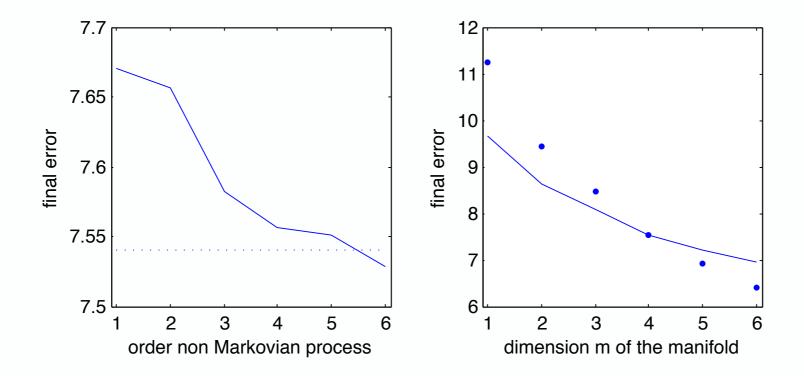


Predictive uncertainty as a function of the dimension m of the reduced dynamical manifold and the order r of the process.



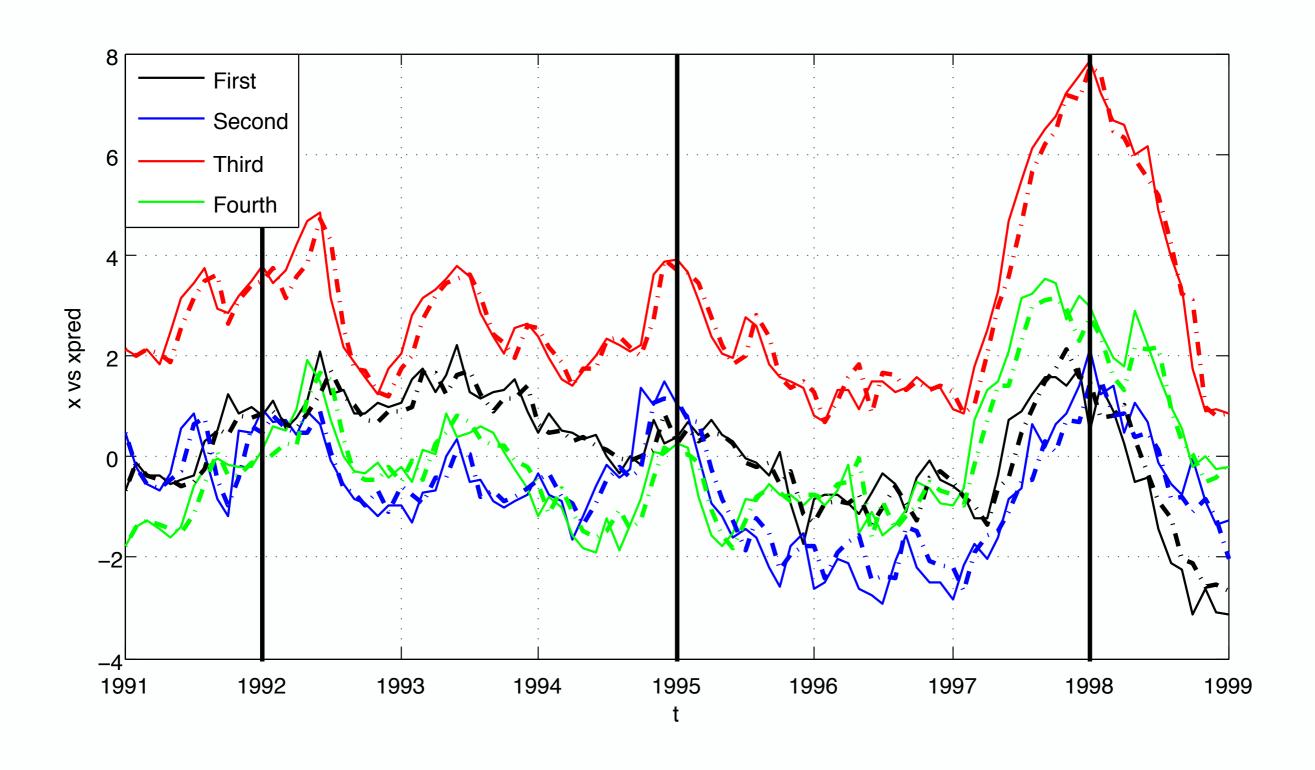


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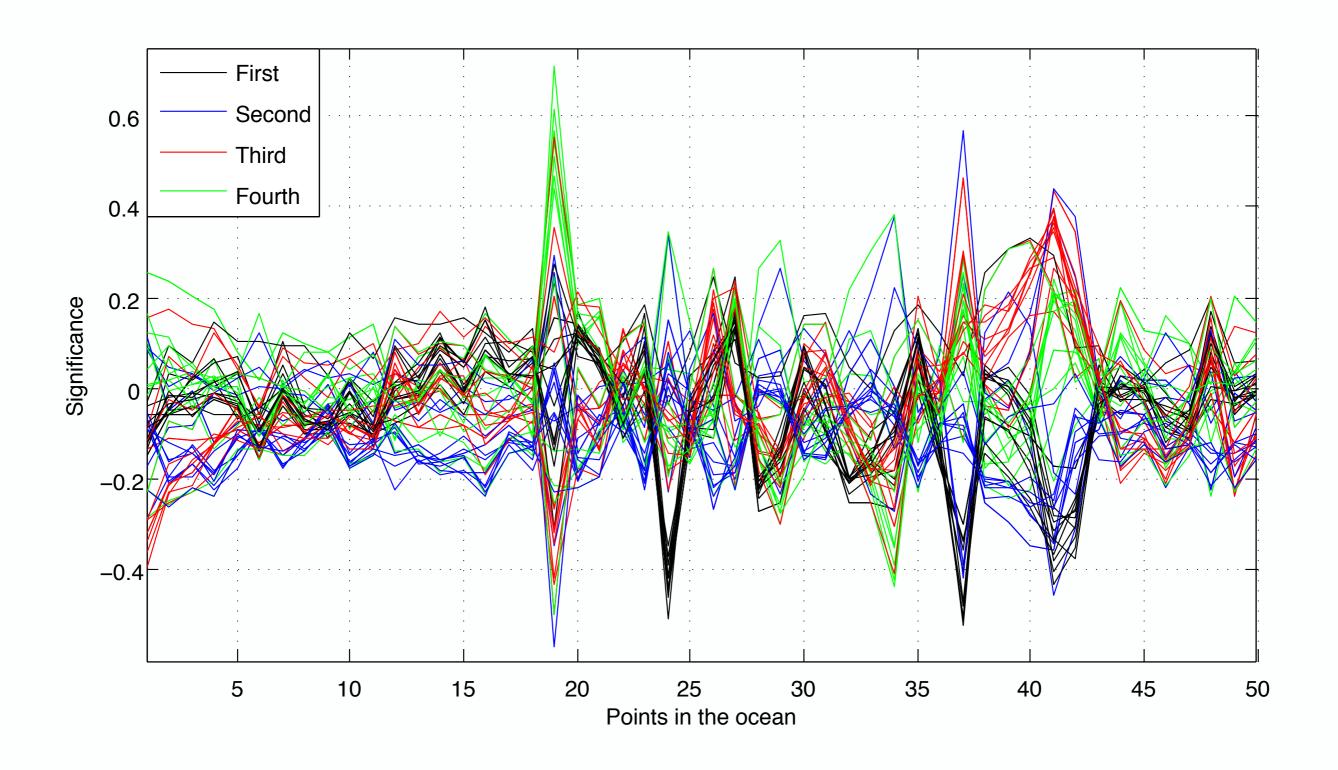


Therefore, we pick r = 3 and m = 4.

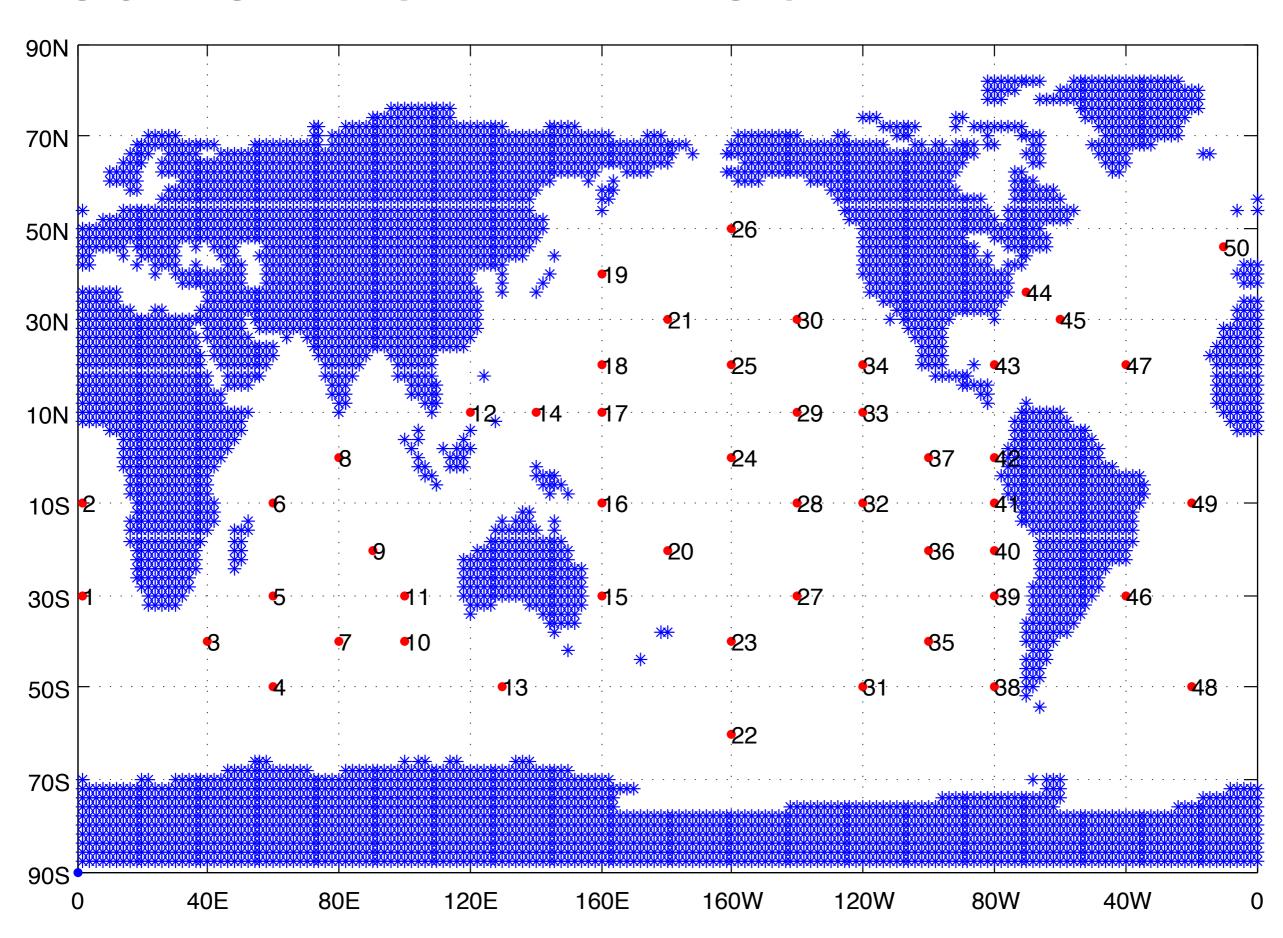
# REAL AND PREDICTED DYNAMICAL COMPONENTS



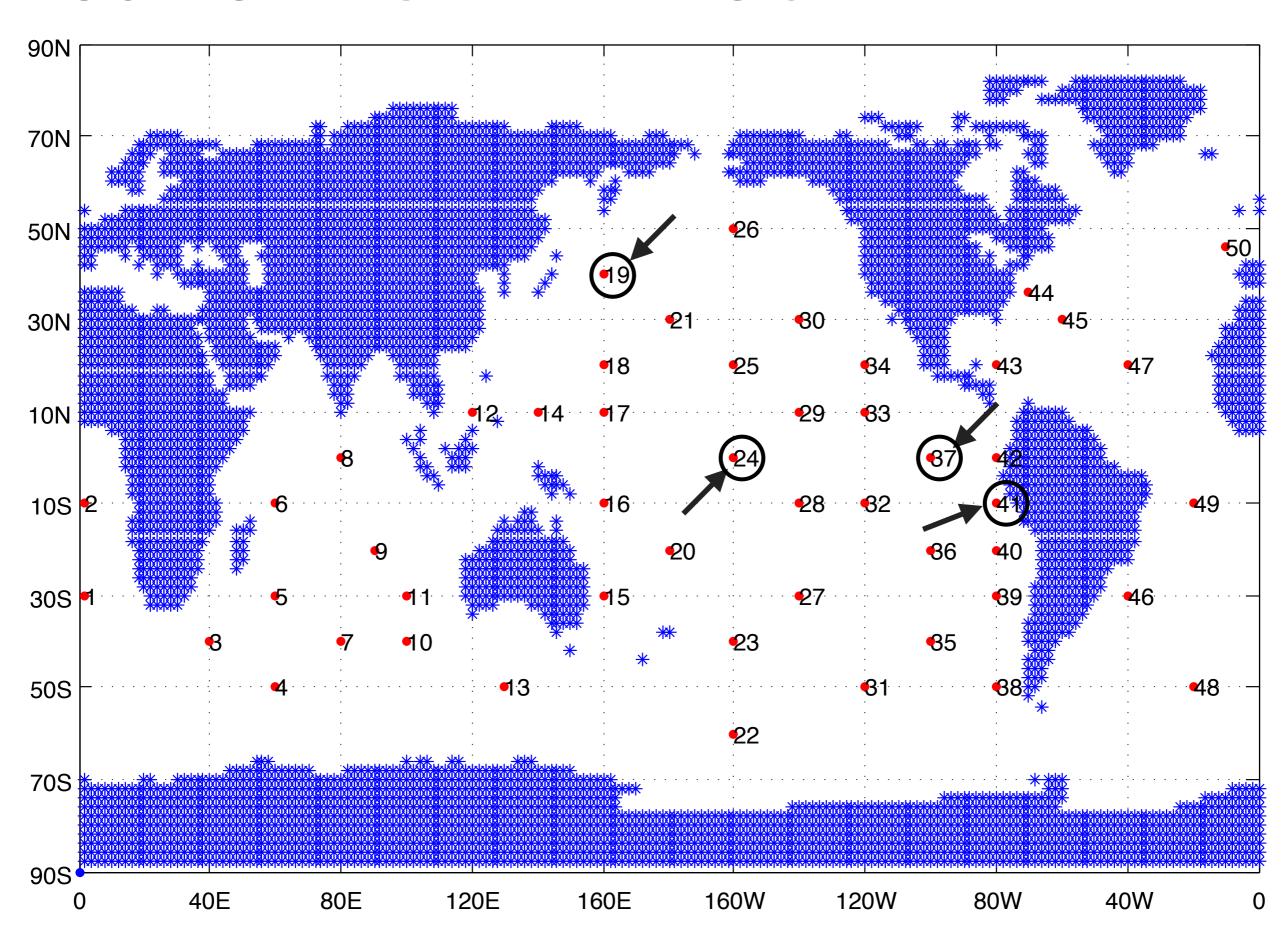
### POINTS OF SIGNIFICANCE: 19, 24, 37, 41



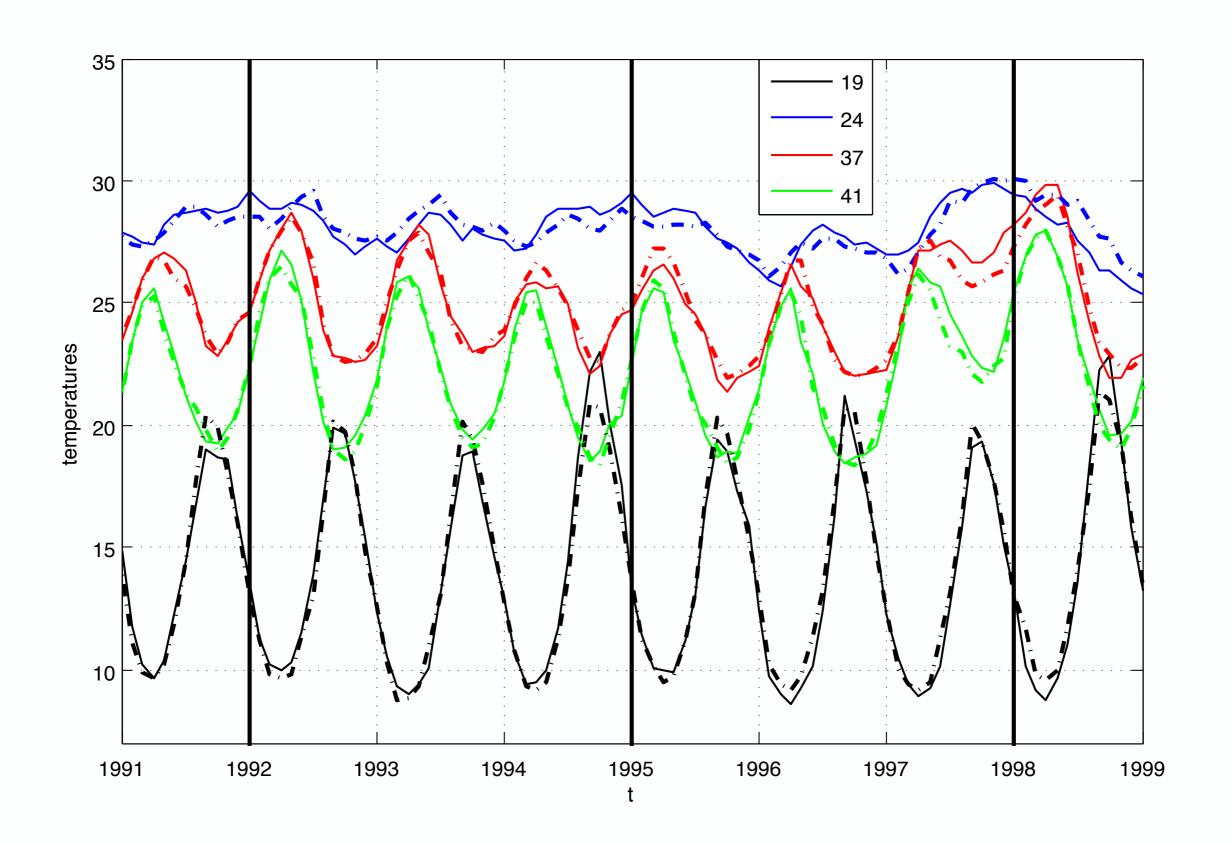
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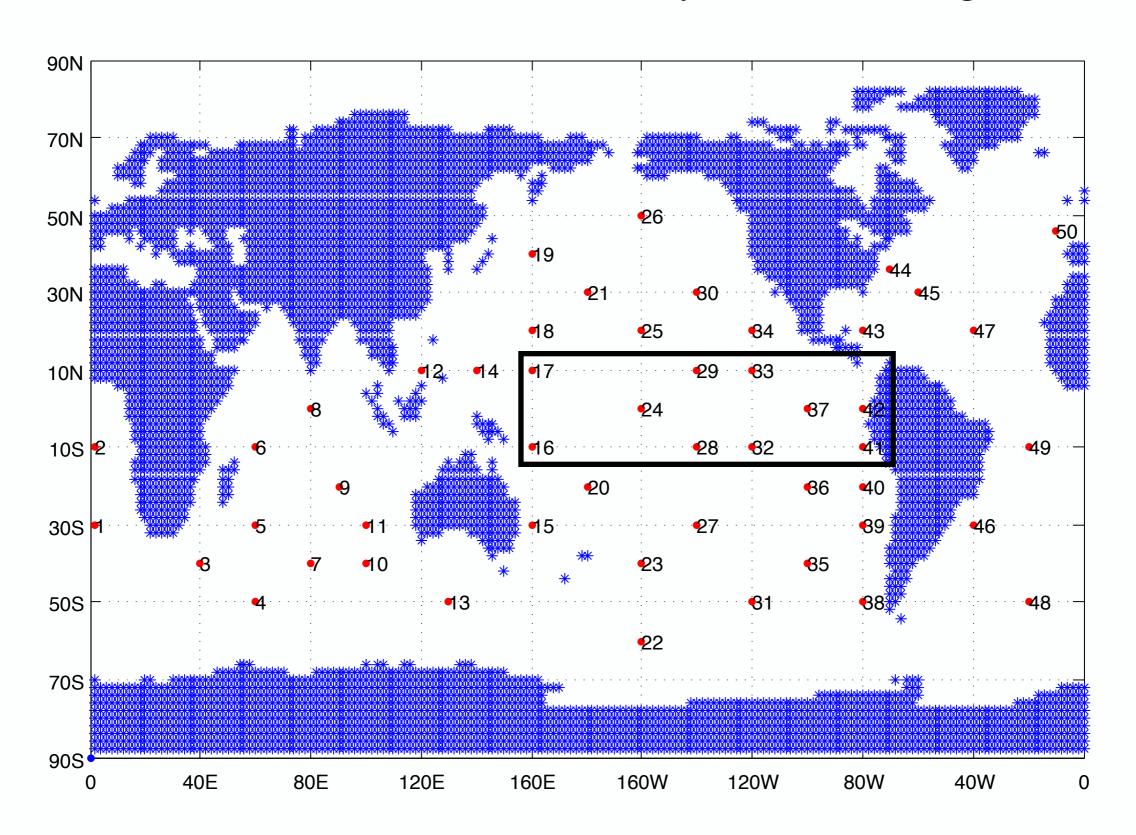


#### OBSERVED AND PREDICTED TEMPERATURES

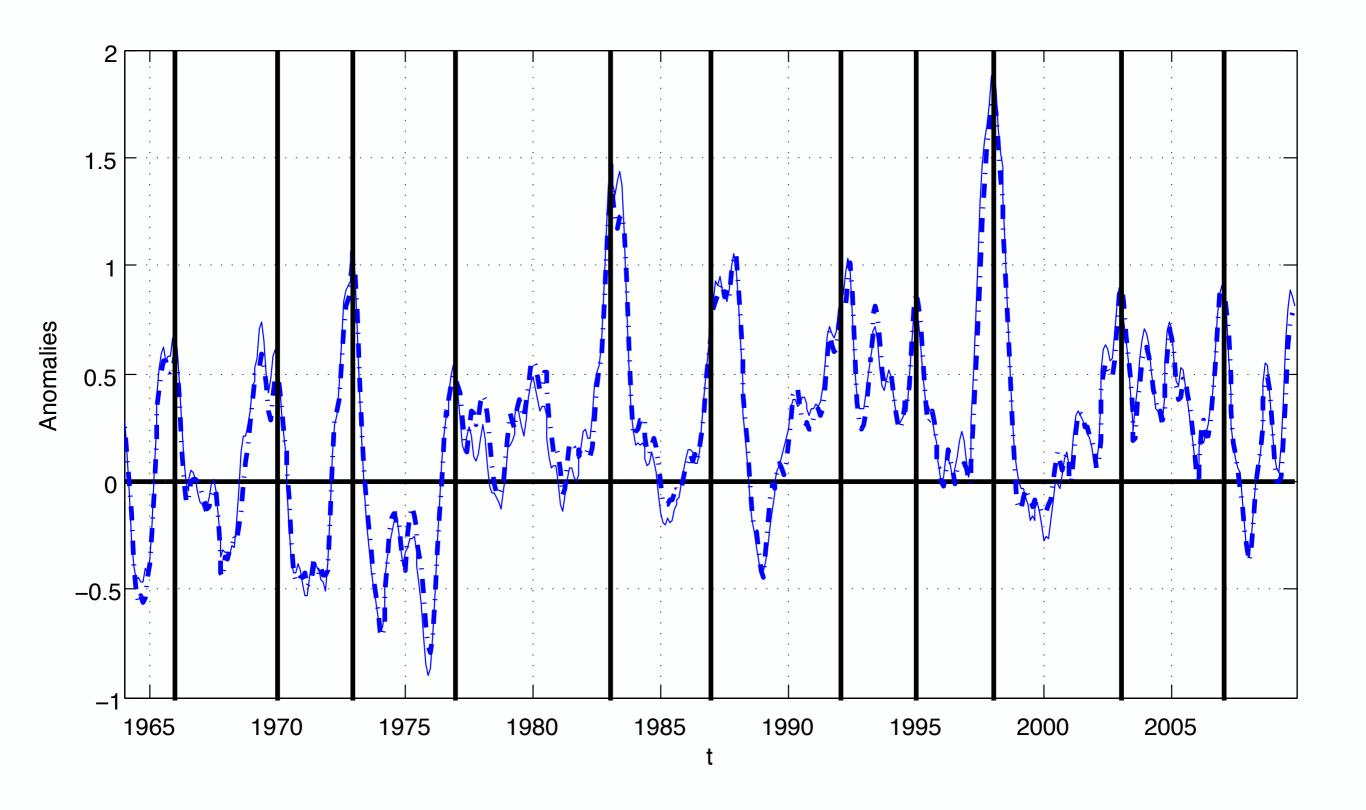


#### ANOMALIES

We consider three-month mean SST anomaly in the following El Niño region:



### ANOMALIES



• This new methodology allows a dimensional reduction of time series seeking a low-dimensional manifold x and a dynamical model  $x_{j+1} = D(x_j, x_{j-1}, \ldots, t)$  that minimize the predictive uncertainty of the series. We have tested on synthetic data and with a real application on sea-surface temperature over the ocean.

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